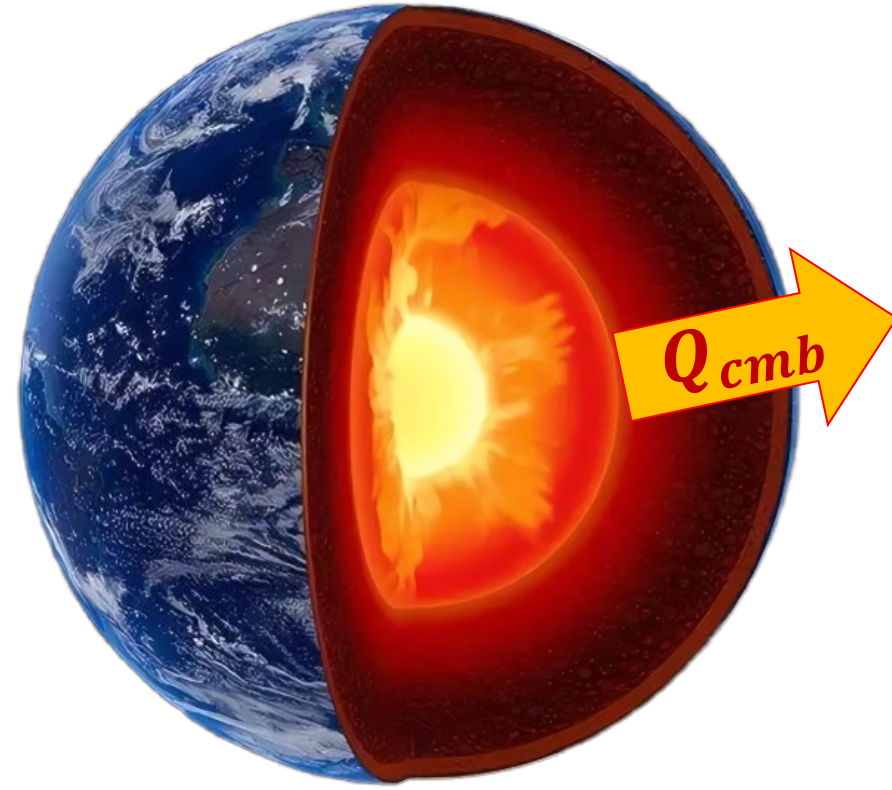


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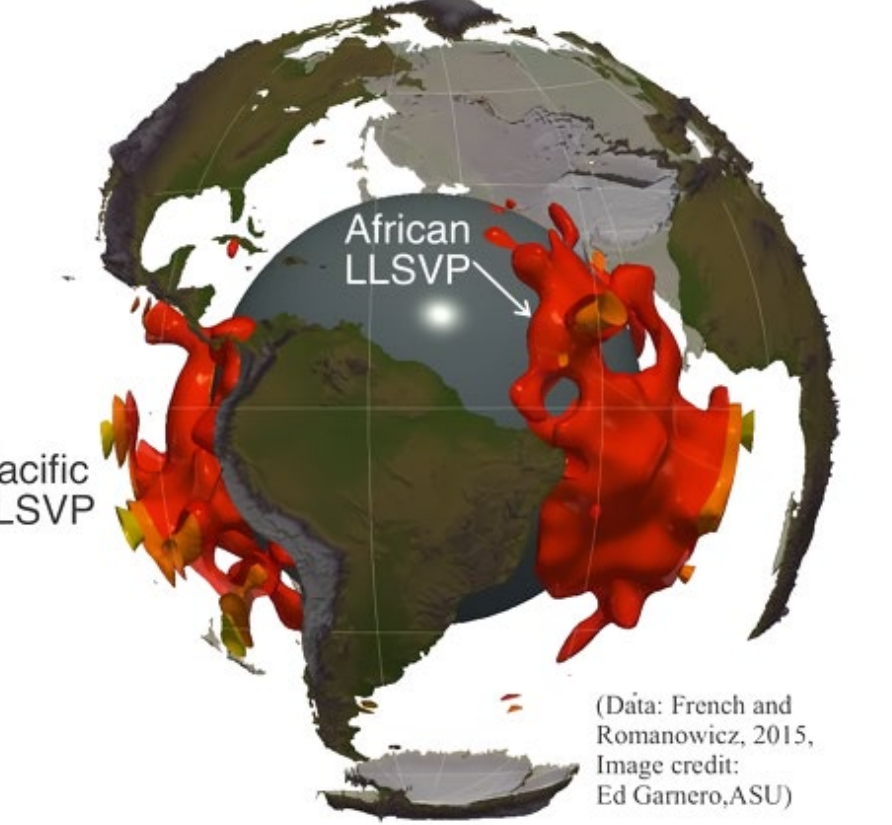
Introduction



The heat flux from the core (Q_{cmb}) is crucial for understanding:

- The growth and structure of the Earth's core
- The intensity and reversal frequency of the geomagnetic field
- Plume size, and so on.

Large low shear velocity provinces (LLSVPs) with compositional and thermal anomalies should have a large impact on Q_{cmb}



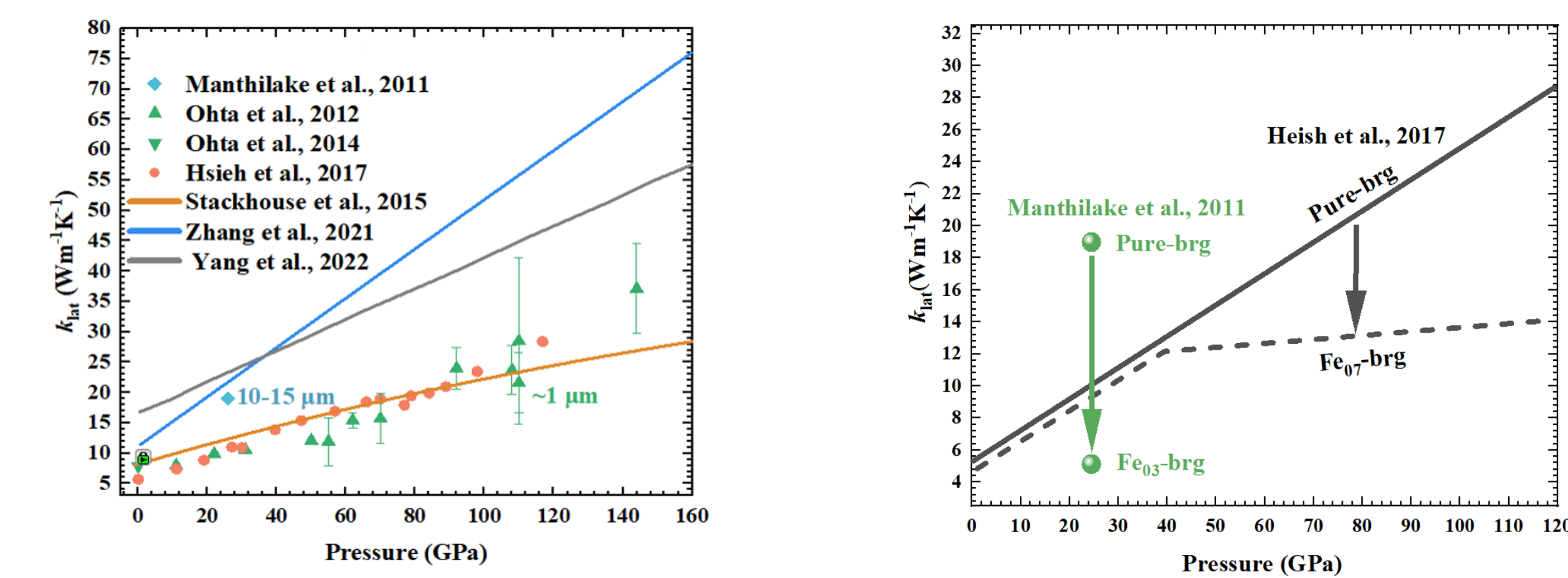
Determining the heat flux in the LLSVPs regions requires the thermal conductivity of the main minerals in the LLSVPs

$$Q = k \frac{\partial T}{\partial x}$$

k : thermal conductivity
 $\frac{\partial T}{\partial x}$: temperature gradients

The thermal conductivity of the potential main minerals (bridgmanite and post-perovskite) in the LLSVPs are highly debated

e.g. Pure bridgmanite (Brg) at 300 K The effect of Fe is also controversial

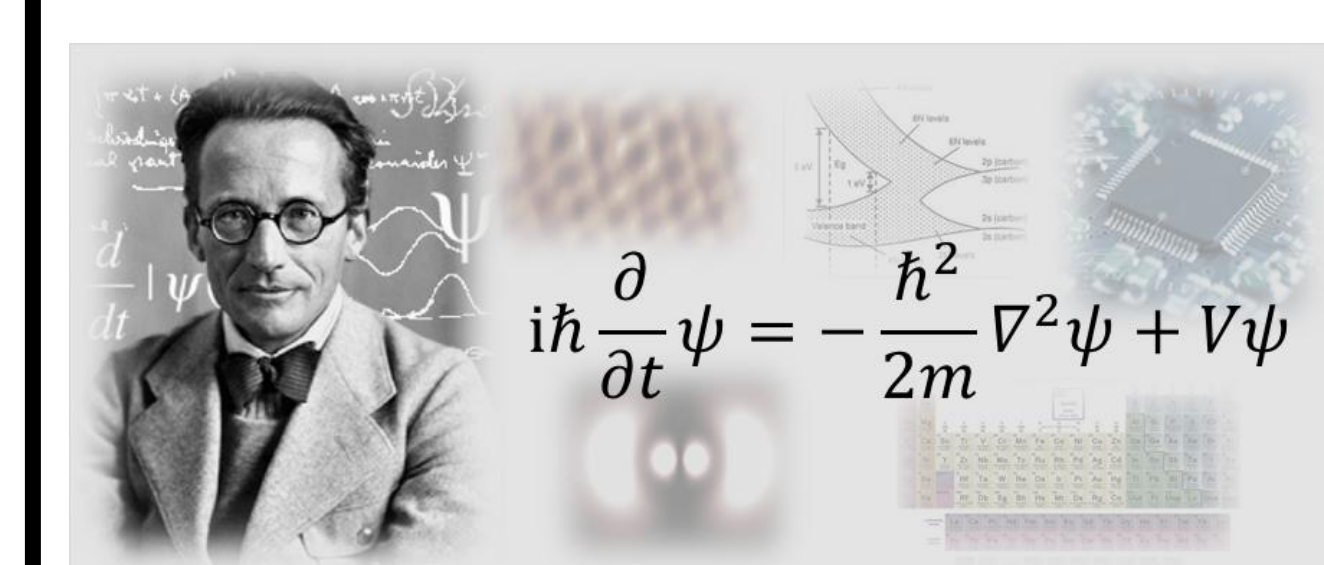


Challenges in determining k at mantle conditions

Experiments: In-situ measurements are challenging
Theoretical calculations: Finite size effect

Impossible for first-principles calculations due to computational costs

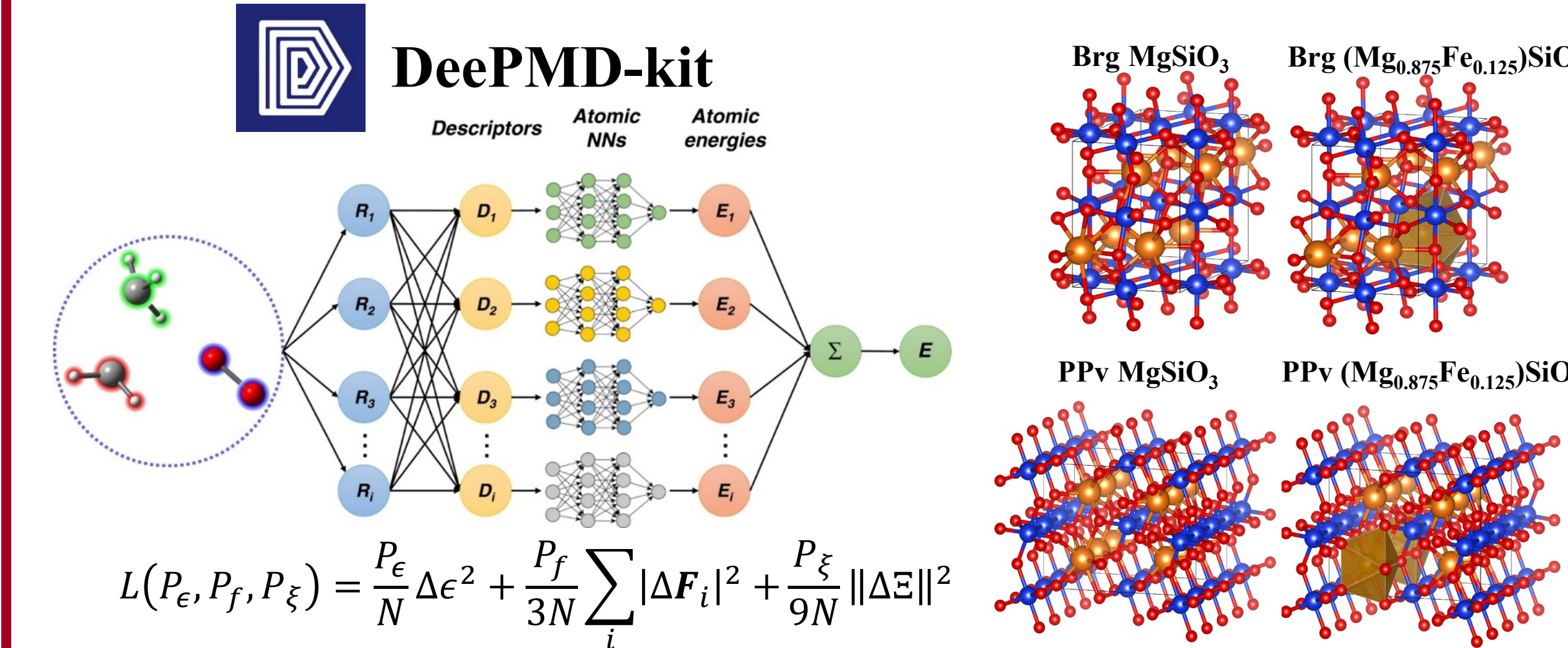
Calculations on huge systems are required



Cost $\propto N^3$, N : number of atoms

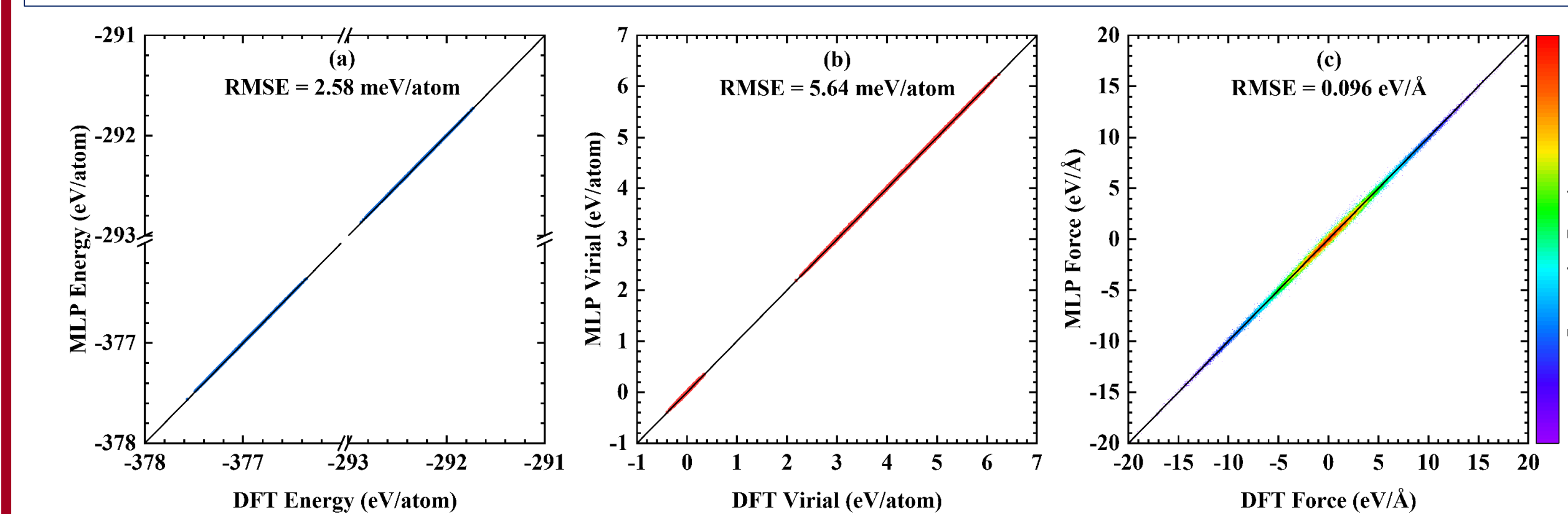
Computational Methods

1, We trained a machine learning potential (MLP) for Fe-free and Fe-bearing bridgmanite (Brg) and post-perovskite (PPv) based on first-principles data



- There are four structures under current consideration. All Fe remains in the high spin state as Si-site Fe was not considered.
- The size of the training set is ~200,000.

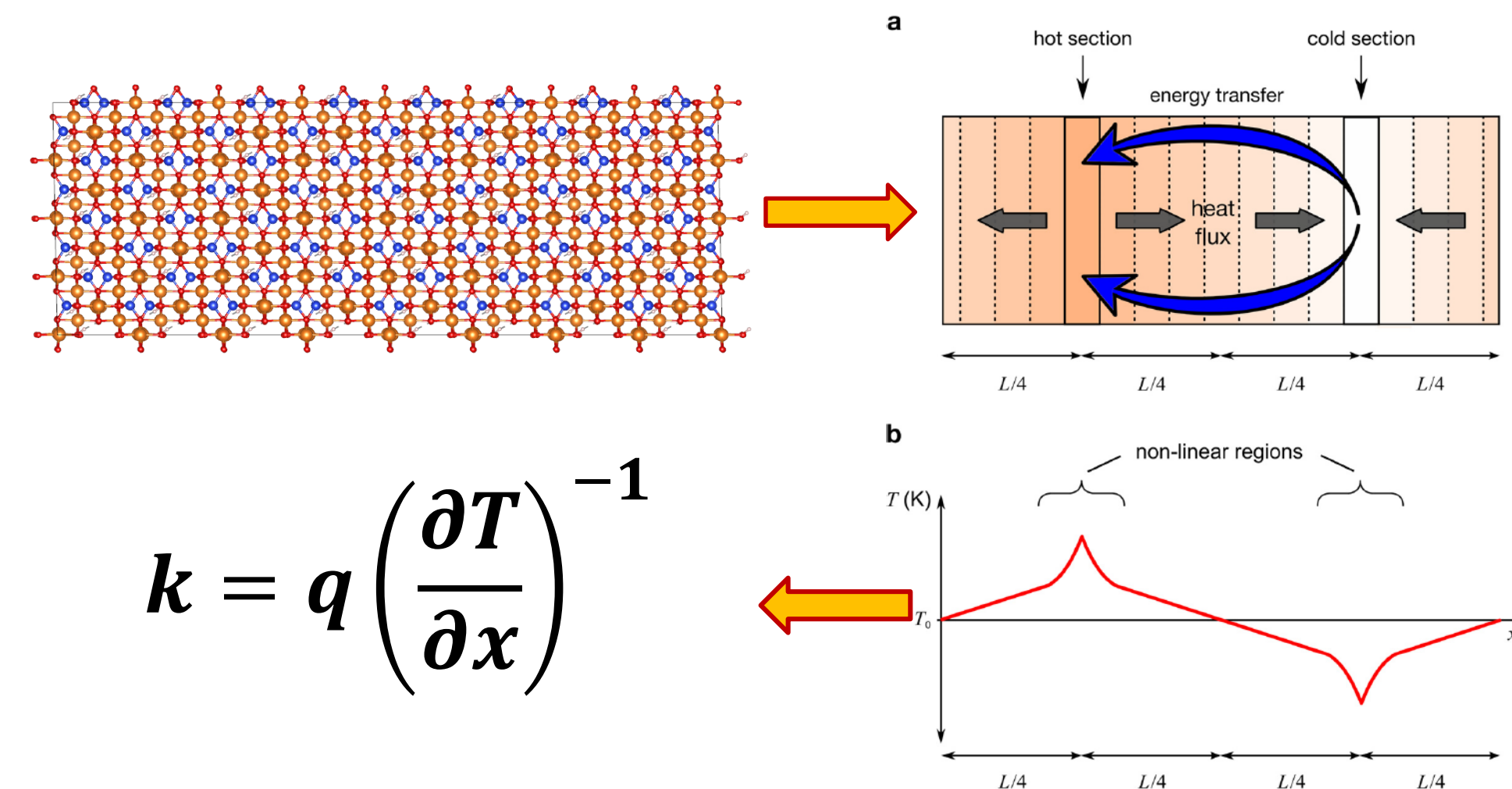
2, The test error of the MLP



- The test set contains ~9,000 randomly selected data.
- The root-mean-square error of energies, forces and virial stresses are 2.58 meV/atom, 0.096 eV/Å, and 5.64 meV/atom, respectively.

3, Using Non-equilibrium molecular dynamics based on the MLP, we predicted the lattice thermal conductivity (k_{lat}) of Fe-free and Fe-bearing Brg and PPv in large cells

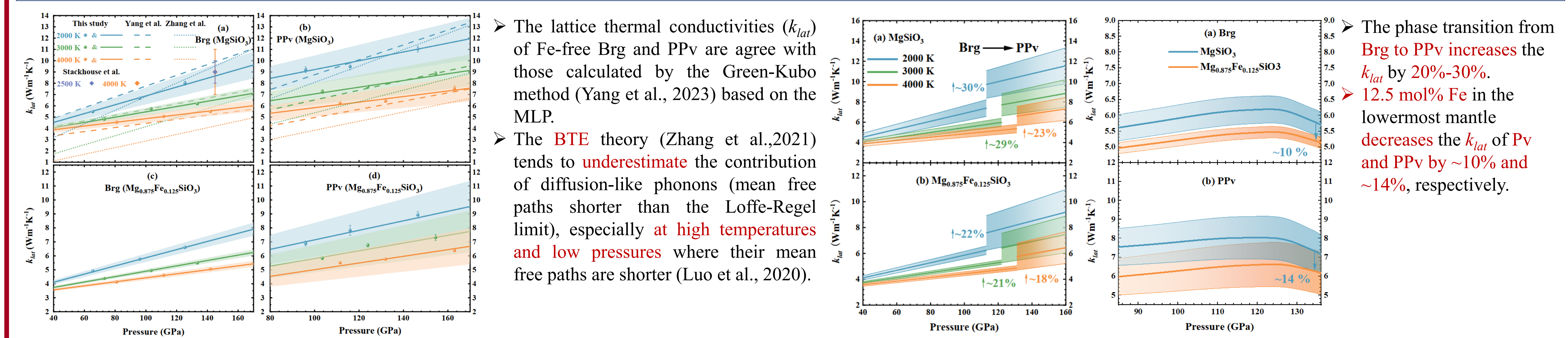
Non-equilibrium molecular dynamics (An experiment-like approach)



- We divided the cell into 32 regions of equal size, and exchanged the velocities of two atoms between the first and seventeenth regions every n steps.
- Over time, a steady heat flow was established and a linear temperature profile was induced. The thermal conductivity can be calculated from Fourier's law.

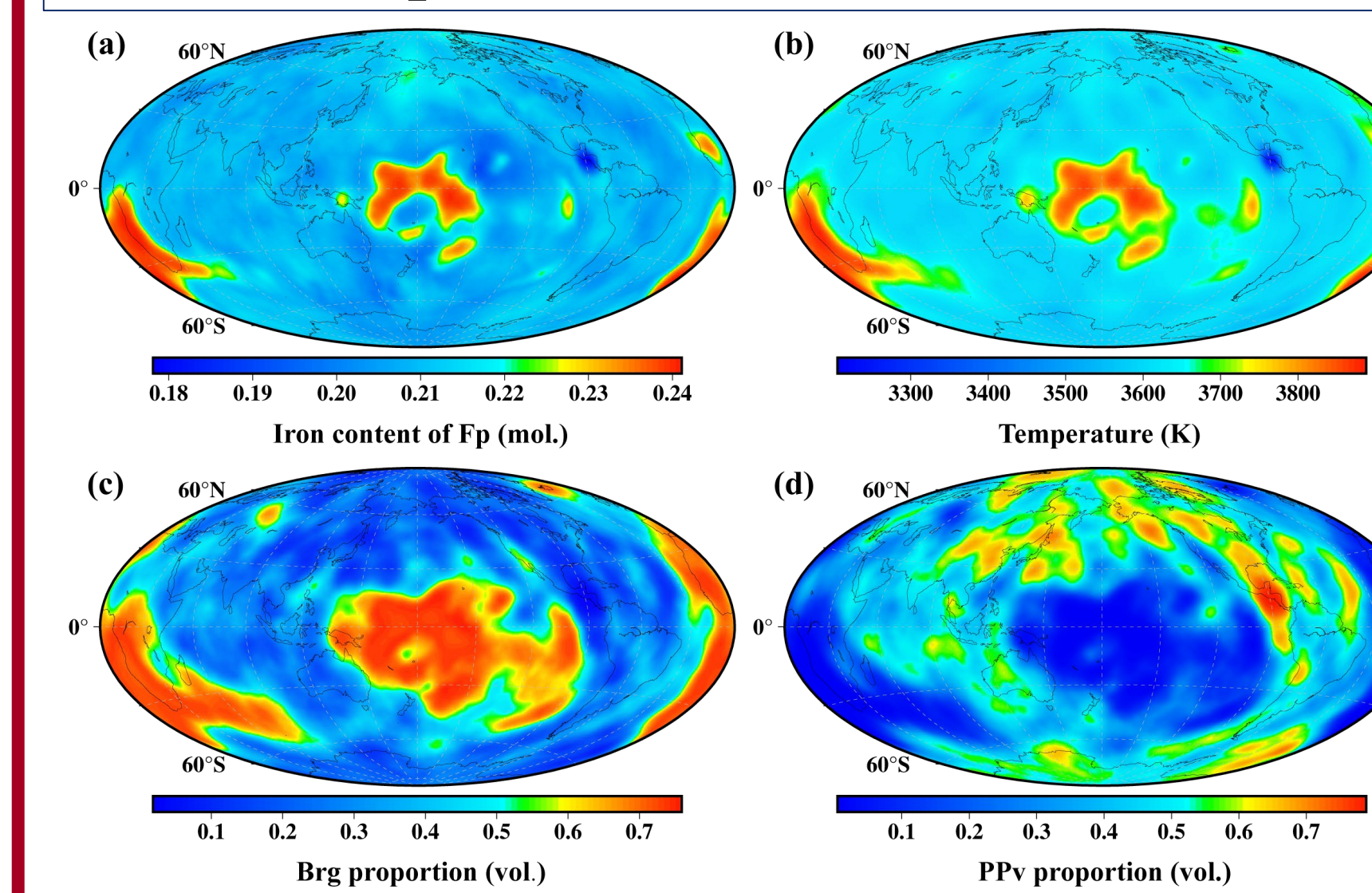
Results and Discussion

The k_{lat} of bridgmanite and post-perovskite



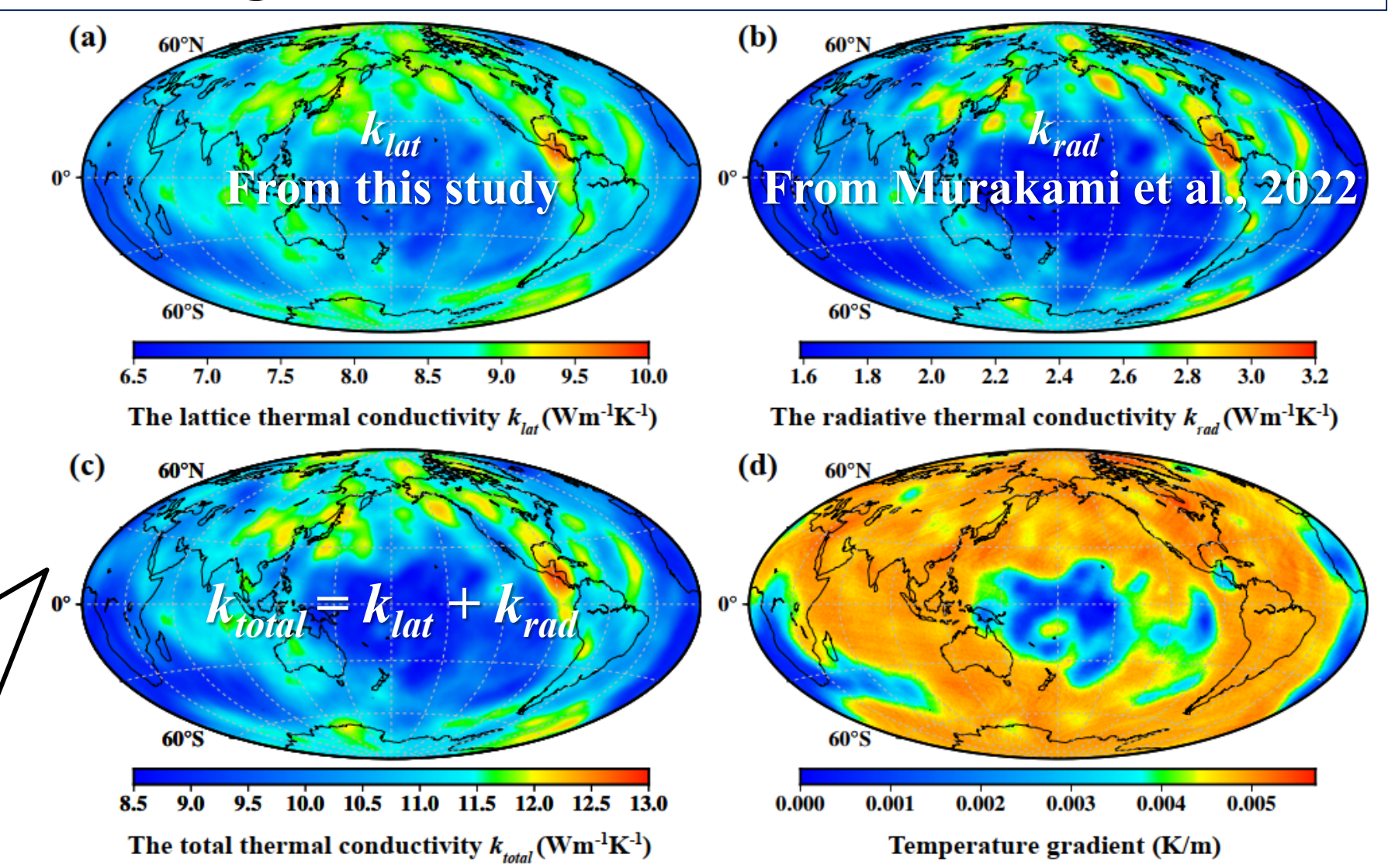
- The lattice thermal conductivities (k_{lat}) of Fe-free Brg and PPv are agree with those calculated by the Green-Kubo method (Yang et al., 2023) based on the MLP.
- The BTE theory (Zhang et al., 2021) tends to underestimate the contribution of diffusion-like phonons (mean free paths shorter than the Loffe-Regel limit), especially at high temperatures and low pressures where their mean free paths are shorter (Luo et al., 2020).

Compositional and thermal state of the lowermost mantle

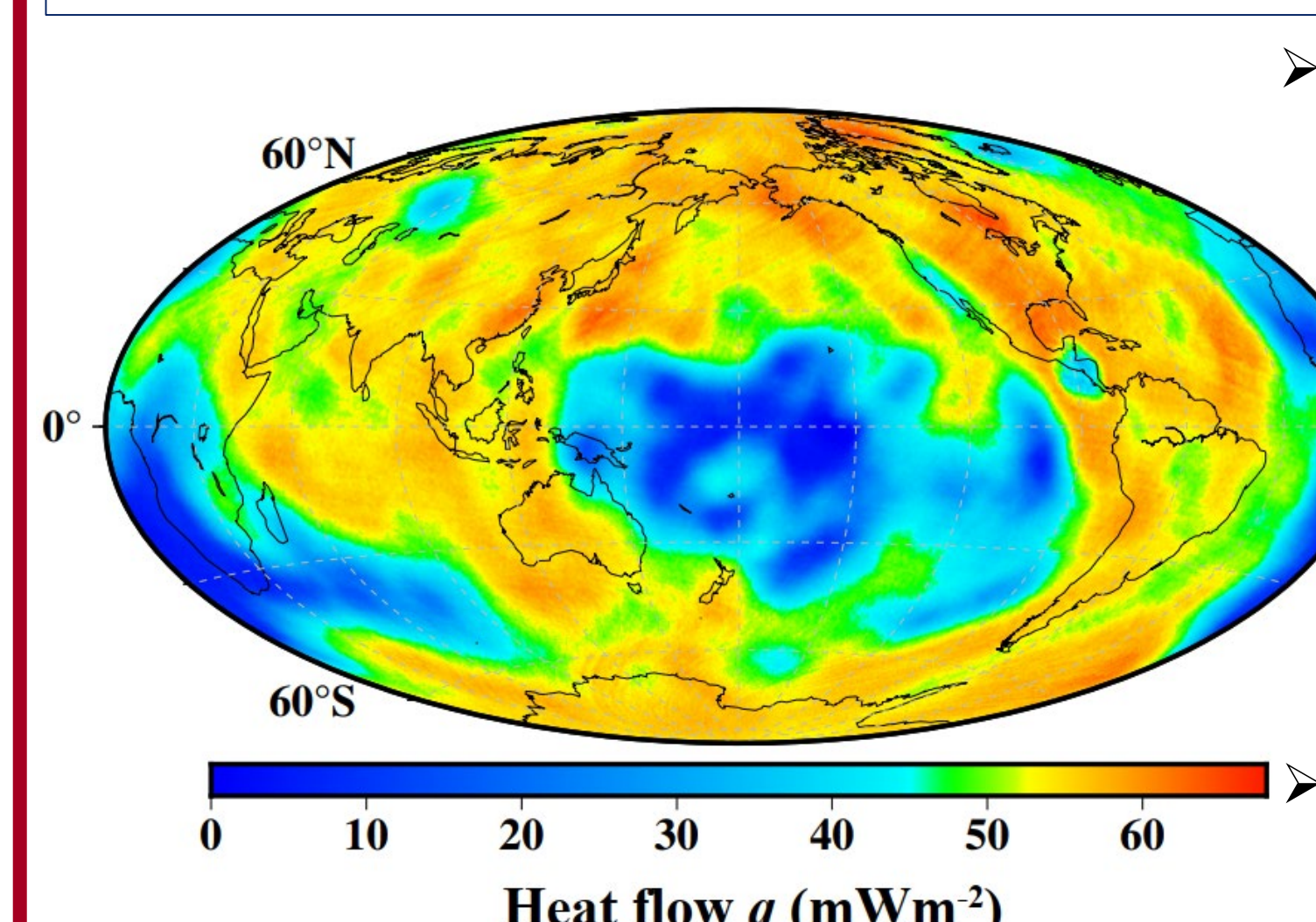


- The inversion based on mineral elasticity and seismic tomography models revealed that the LLSVPs regions are enriched in iron-rich Brg (Deng et al., 2023).
- The thermal conductivity is relatively low in the LLSVPs regions due to higher temperatures and lower PPv content within the LLSVP regions.
- The temperature gradients are relatively low in the LLSVPs regions due to the high temperature there.
- The k_{lat} of davemaioite and ferropiclasite are from previous studies (Zhang et al., 2021; Hsieh et al., 2018; Mantihalake et al., 2011).

The k and temperature gradient at the lowermost mantle

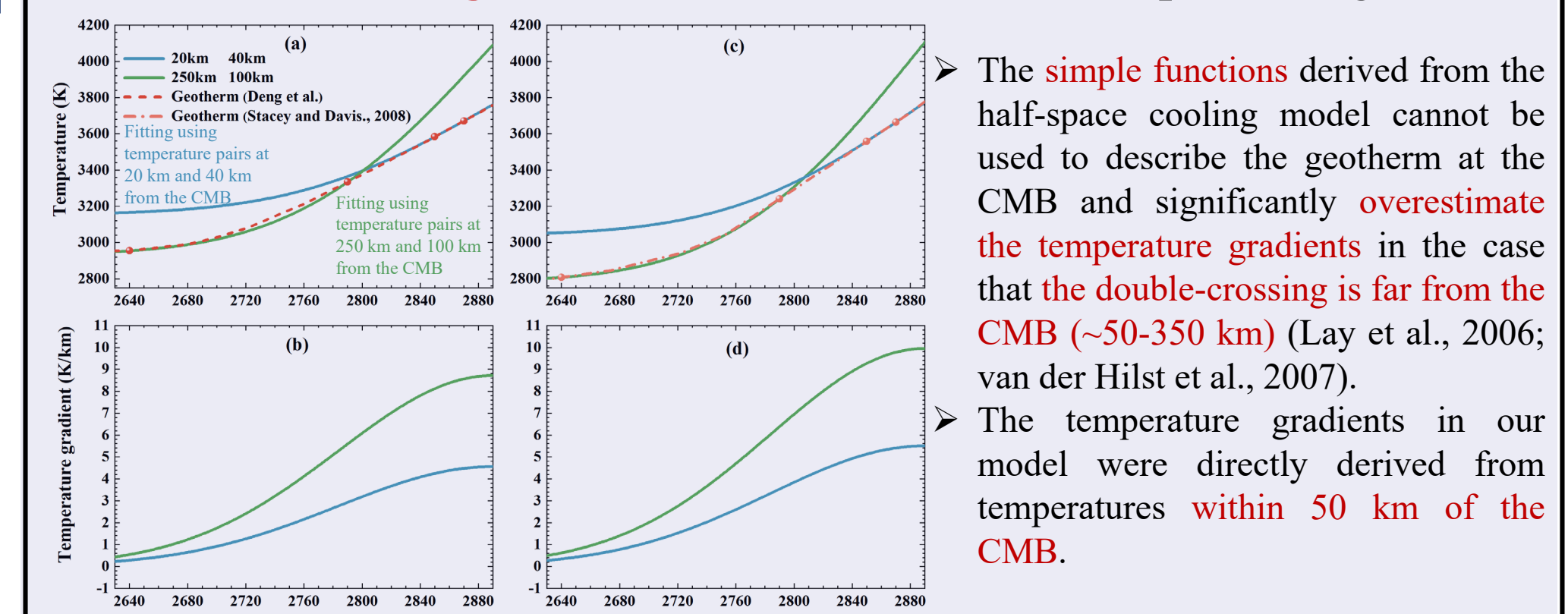


The heat flux from the Earth's core



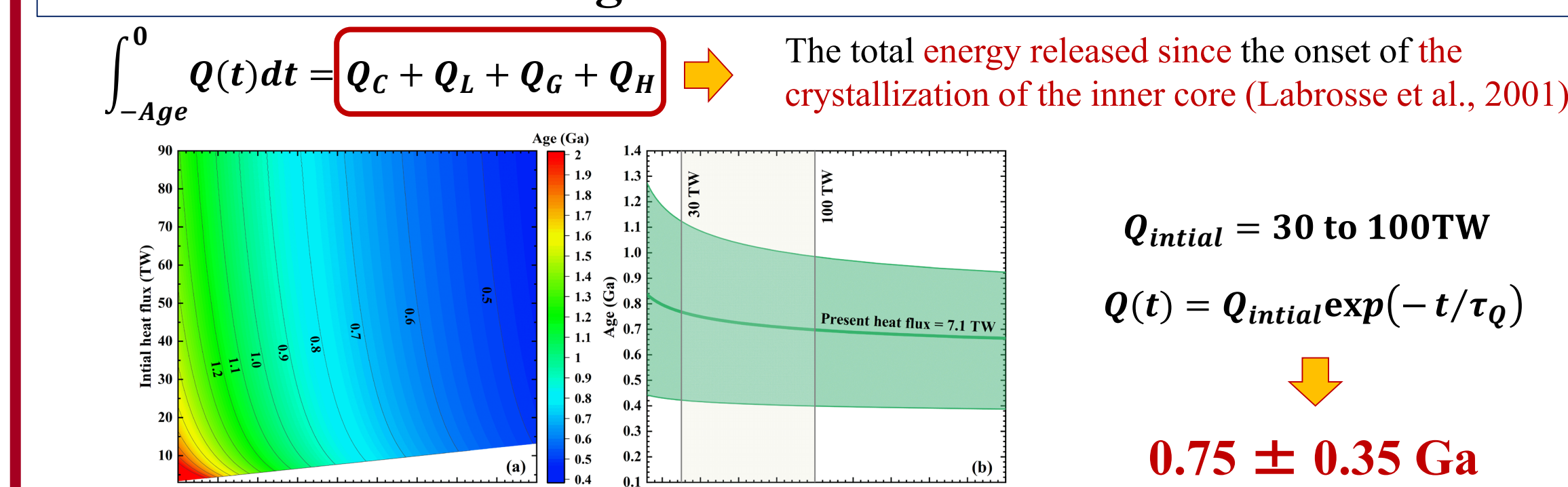
- The heat flux in the LLSVPs regions is extremely low due to the lower temperature gradient and thermal conductivity there, which leads to large lateral variations ($q^* = \frac{Q_{max} - Q_{min}}{2 \times Q_{average}} = 1$).
- Thermal convection in the Earth's core below the LLSVPs regions is suppressed by the lower heat flux, which affects the geomagnetic field.
- Total heat flow from the Earth's core is 7.1 ± 0.5 TW, which is sufficient to sustain geodynamo (Nimmo, 2015), but lower than previous estimates from the double crossing of the Brg-PPv phase boundary (Lay et al., 2006; van der Hilst et al., 2007).

The double crossing method overestimates the temperature gradients



	Thermal conductivity	Temperature gradients	Total heat flow
This study	10.4 W/mK	4.4 K/km	7.1 ± 0.5 TW
Double crossing	10 W/mK	6-16 K/km	12 ± 5 TW

Age of the inner core



- We trained an MLP for Fe-free and Fe-bearing Brg and PPv with data from first-principles calculations, and then investigated their lattice thermal conductivity at mantle conditions in the large cells with finite-size effects well considered.
- The phase transition from Brg to PPv increases the lattice thermal conductivity by 20%-30%.
- Above the CMB, the presence of 12.5 mol% Fe decreases the thermal conductivity of Brg and PPv by ~10% and ~14%, respectively.
- The LLSVPs regions with extremely low heat flux hinder the cooling and thermal convection of the core.
- The total heat flow from the Earth's core was estimated to be 7.1 ± 0.5 TW, which supports a geologically young inner core of 0.75 ± 0.35 Ga.

Summary

